

Sediment content of nearshore fast ice - Fall 1980, Beaufort Sea, Alaska

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INTRODUCTION

The sediment content of the seasonal fast ice off the north coast of Alaska has been shown to be a significant part of the sediment transport system for fine-grained materials (Barnes et al. , 1982). Sediments are believed to be incorporated into the **ice** canopy during fall storms. In the fall **of** 1980 only minor wind and wave events occurred during the initial freezing of the fast-ice canopy. A sampling program was carried out to provide a comparison of the sediment load carried by the fast **ice** in a season of initial storms with 1977 when major wind and wave events are known to have existed during freezeup (**Barnes** et al., 1981). We report the results of this sampling effort in this report.

The details of the ice regime and the history **of** sediment studies in the ice canopy **have** been outlined by Barnes and others, **1982**, and will not be presented here. The unique aspects of shelf ice distribution in the fall of 1980 when the ice sampled for this report formed, are 10-20 km of open water over the inner shelf. Off the central Beaufort Sea the pack ice was offshore during the early fall and freezeup. Coastal winds were moderate northeasterlies with peak sustained velocities measured at 20-30 mph (Fig. 2). Freeze-up occurred over **a** 5-7-day period and was followed by continued moderate wind events (Fig. 2).

METHODS

The area of study extends from Cape Halkett on the west to **Flaxman** Island on the east (Fig. 1). All cores were taken from the fast ice using a 5-cm coring device. Samples from the eastern half of the study area extend out to, but do not include, the **stamukhi** zone or the pack ice seaward of the **stamukhi** zone. Cores were taken in areas of flat-lying ice **away** from the small ridges and hummocks. Ridging and rafting of small ice blocks **was** common during freeze-up, mechanically thickening the ice canopy. Our coring program in areas **away** from these features represents ice thickness due to normal ice growth and does not represent ice thicknesses due to rafting and ridging. Furthermore, our cores thus represent an average ice thickness less than is actually present in the fast-ice zone. As the season progresses and ice thickness increases to incorporate many of the smallest ridges and hummocks, the regional variations in the thickness of the fast ice become less and less.

The methods used to obtain and analyze the cores used in this study essentially duplicated those used in our earlier study (Barnes et al., 1982). Cores were returned to the laboratory in the frozen state and photographed **on** a light table to accentuate stratigraphy. The cores were **subsampled**, allowed to melt, and then filtered according to the **techniques** of Drake et al., (1972). Salinities were determined using the index of refraction method.

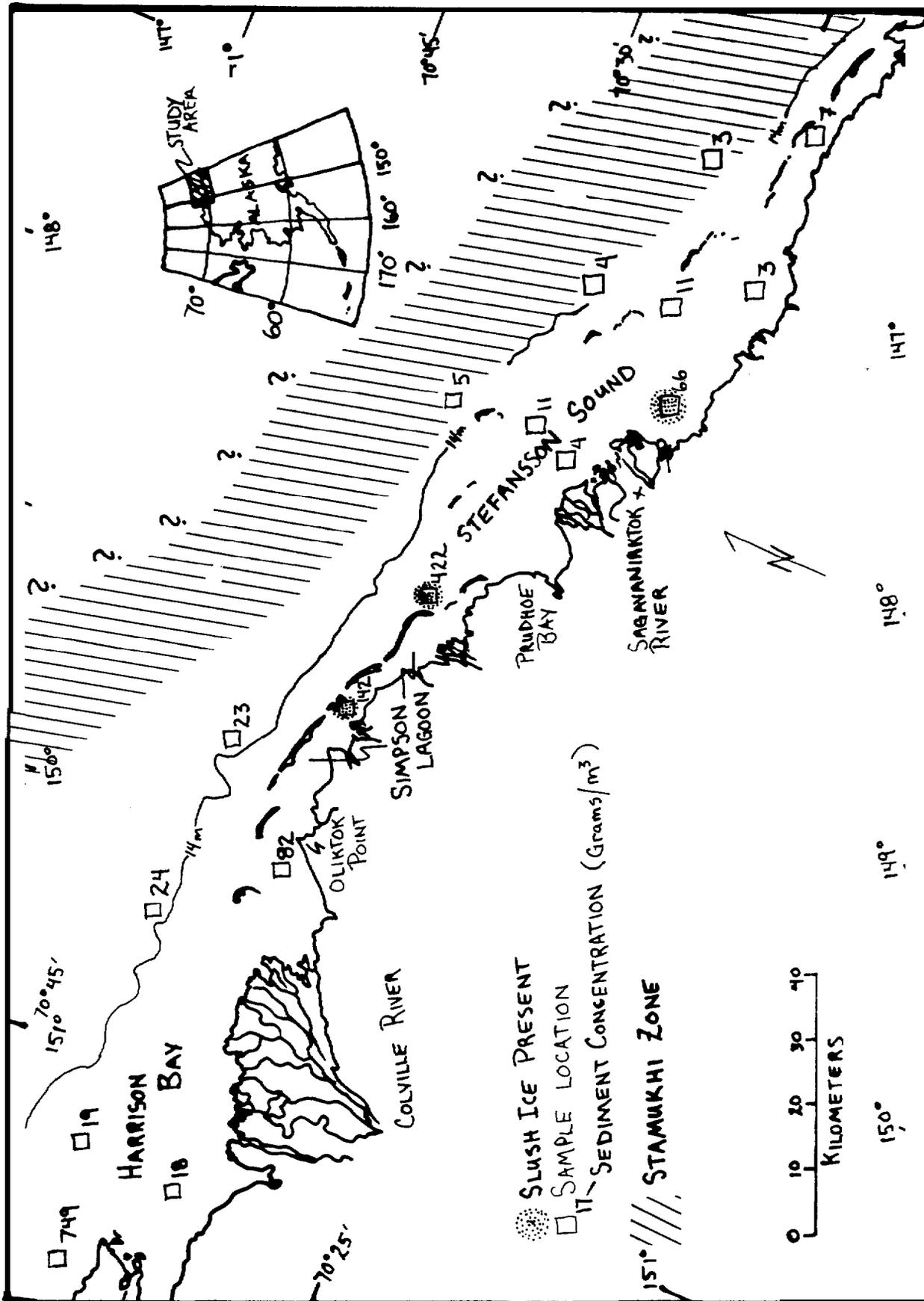


Figure 1. Sediment concentrations at ice core sites sampled in November, 1980. Fast-ice zone is the ice area inside the stamuksi zone.

OBSERVATIONS

Stratigraphy

The cores and samples taken from the fast ice in 1978 had a regional **stratigraphic** similarity (Barnes et al., 1982) which was not found *in* the cores taken during the fall of 1980. As in 1978, the surface 5-10 **cm of most** 1980 ice cores appear less sediment-laden than the underlying sections of the cores. Below this surface layer the general character was **highly** variable-- some cores being extremely sediment laden and others essentially **sediment-free**. The location of the sediment-laden zones varied from upper and middle parts of the cores to the very base of the cores where new ice was still being formed (see appendix). Salinities of cores decreased with depth in core. One core of the **Sagavanirktok** River was taken in fresh-water ice.

Sediment Concentrations

Sediment concentrations from individual ice core segments varied from less than 3 **mg/l** to a high value of 2127 **mg/l** (see appendix). In order to compute the quantity of sediment in the fast ice, the segment concentrations in each core segment were normalized to a value for the entire ice thickness (core length) at each sample site. When this was done the sediment concentration values ranged from 3 **mg/l** to 749 **mg/l** (Table I) over core lengths of 35-55 cm. The **areal** distribution of sediment suggests higher concentrations inshore and on northeast-facing coasts (Fig. 1). Lowest values occur both inshore - as off the Sagavanirktok delta - and offshore in the vicinity of the stamukhi zone north and east of **Prudhoe** Bay.

While drilling at 3 inshore sampling sites, a sediment-laden soft (**slushy**)ice mass of unknown thickness was encountered below the core (Fig. 1). The concentration and amount of sediment in this ice mass is not known.

Sediment Textures

Binocular-microscope investigation of filtered sediment indicates that **silt-** and clay-sized particles make up the vast bulk of sediments in the cores. When sand-sized material was present it often was in the form of fibrous organic matter rather than mineral grains. The core with the largest sand fraction (estimated at about 10 percent by weight) occurred in the core from Simpson Lagoon (Fig. 1).

Weather during freeze-up

During the **fall** of 1980 freeze-up - that period when the sea goes from ice-free to essentially ice-covered - occurred between the 20th and 24th of September (E. Reimnitz, personal **commun.**). During this time abundant quantities of **frazil** ice were present along the outer coasts of the islands and ^{were} noted to be accumulating against coastal promontories such as **Oliktok** Point, the bend in Long Island, and in the vicinity of artificial islands in Stefansson Sound.

Table I - Summary of Ice Core Characteristics

Core	Average Sediment Concentration (gm/m ³)	Core Length (cm)	Salinity (‰/00	Notes
2	749	37.5	2 - 4°/00	
3	19*4	35*5	2 - 4 ⁰ /00	
4	18.1	42.0	0 - 4°/00	
5	24.0	47*5	2 - 6°/00	
6	82.2	39.0	2 - 3 ⁰ /00	
7	22.8	51.0	2 - 4°/00	
8	141.8	55.0	0 - 2°/00	
9	421.6	54.0	0 - 4°/00	- salty at surface fresh at base
10	3.5	46.0	0 ⁰ /00	- fresh water
11	5.4	53.0	1 - 4 ⁰ /00	
12	10.5	41.0	2 - 3°/00	
13	66.1	53.5	0°/00	- fresh water(?)
14	3.6	44.5	0 - 2°/00	
15	10.8	39.0	0 - 4°/00	
16	3.3	37.5	0 - 5°/00	
17	3.2	41.5	1 4°/00	
18	<u>7.2</u>	<u>38.5</u>	<u>2 - 5°/00</u>	
Ave ra ge	97.8	44*5		

Coastal wind speed noted at the weather station at Barter Island 100 km to the east of the study area where low-velocity northeasterlies 5-10 mph prior to freeze-up rising to about average winds of 20 mph during the first part of freeze-up followed by lower velocity westerlies at the end of freeze-up (Fig. 2).

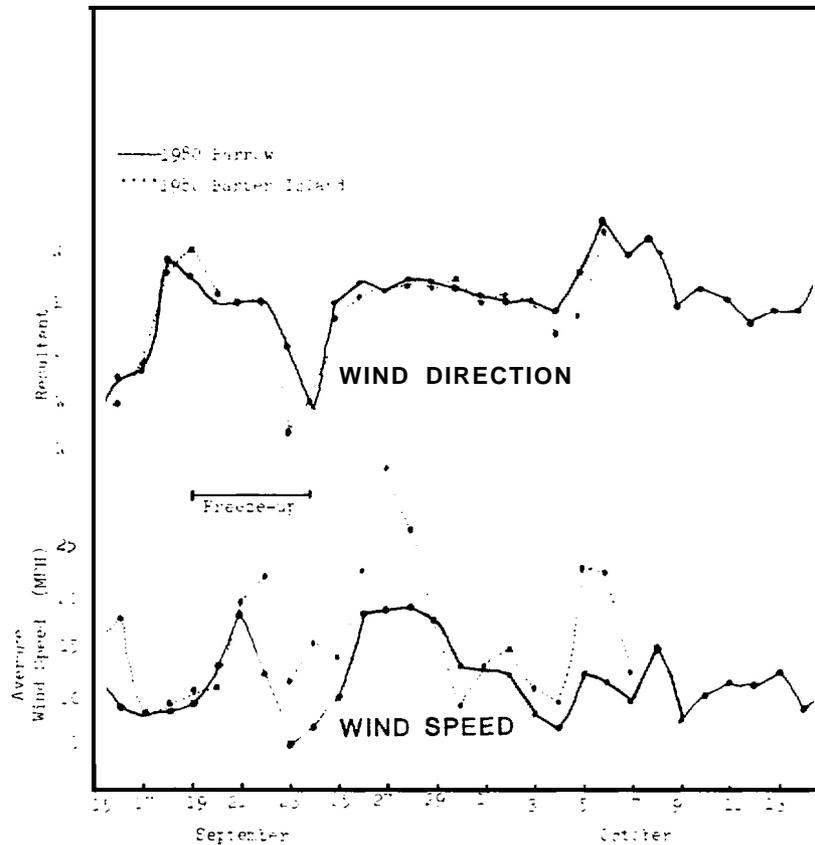


Figure 2. Coastal wind speed and direction during and subsequent to freeze-up, fall 1980.

DISCUSSION AND CONCLUSIONS

When compared to observations of the 1977-78 ice canopy, sediment concentrations in the 1980-81 are about an order of magnitude lower (Table II, Barnes et al., 1982). We attribute these lower values to the less intense wind and wave regime during the fall of 1980 and the fact that the entire freeze-up process occurred over a shorter period of time, thereby preventing any extensive reworking and resuspension of sediments in the presence of frazil ice.

Based on observations in this study we believe that the mechanism suggested by Barnes et al. (1982) is basically valid. Their mechanism for incorporating sediments into the fast-ice canopy calls for suspension of frazil ice and sediment to be created during freeze-up. Accumulations of the sediment and ice slush would occur on the windward side of promontories.

Extensive areas of open water allow waves to build up, resulting in more intensive sediment resuspension. In 1977 open-water areas were more extensive and waves were larger than in 1980 which resulted in a larger quantity of sediments in the 1978-79 ice canopy. A lingering freeze-up with prolonged periods of frazil generation, would enhance the quantity of sediments incorporated in the fast ice.

Table II - Sediment in Ice Canopy

Ice Canopy -	No. of cores	Average Sediment Concentration (range) (g/m ³)	Average core length (m)	Area (km ²)	Weight of sediment in ice canopy (t.)
Harrison Bay	4	203 (18-749)	0.41	2140	1.78 x 10 ⁵
Stefansson Sound	7	75 (3-422)	0.44	1300	0.43 x 10 ⁵
Simpson Lagoon	2	112 (82-142)	0.47	146	0.08 x 10 ⁵
Study Area	17	97.8	0.445	9300	4*05 x 10⁵
<hr/>					
Ice canopy - 1977-78					
Study Area	18	243 t/km ²		11,400	28.0 x 10 ⁵
Simpson Lagoon	3	480	1.58	146	1.11 x 10 ⁵

The data for this report also support the observation that the sediments placed in the fast-ice canopy are primarily well sorted silt and clay fractions. The sorting of fine-grained materials is believed to be related to 3 factors. First, the probability is higher that **fine-grained** materials will remain in turbulent suspension for longer periods due to their slower settling velocity and therefore silts and clays **would** be more readily available for inclusion in the ice canopy. Secondly, turbulent energy decreases as **frazil** ice masses become larger and larger thus damping wave energy. In this decreasing turbulence the coarser material will settle out first, thus **fine-grained** materials are most likely to be included in the congealing **frazil** masses. Lastly, it is likely that **frazil** ice in a turbulent water column **impacts** the bottom eroding seabed materials and preferentially driving the more easily suspended fine-grained materials into suspension.

The areas with high concentrations of sediment were less widespread in 1980 than in 1978-79. The data of this report show high concentrations along the coast, and in the shallow lagoons in 1980 while in the 1977-78 ice canopy high concentrations were present to the seaward limit of the study area. This

is believed to be related to the lack of intense storm activity during **freeze-up** in 1980 and a relatively shorter period for the freeze-up. The shorter interval does not allow for sediment-laden masses of the ice canopy to advect to other areas of the shelf prior to becoming stabilized in the ice canopy.

The stratigraphy of the **cores** generally follows that observed in the 1978 cores where the uppermost parts of the cores are reasonably sediment free. Sediment content increases irregularly with depth. In some cases the increasing of sediment-rich zones extends to the bottom of the cores (see appendix) as if all of the sediment-laden ice had not as **yet** been incorporated in the thin-ice canopy (Reimnitz and Dunton, 1979) In support of the above hypothesis, at three of these sites where sediments extended to the bottom of the core (Fig. 1 and appendix) and 13) **we** encountered sediment-laden **frazil** ice at the bottom of the drill hole.

CONCLUSIONS

Ice-core sampling of the fast-ice canopy in the fall of 1980 showed the widespread presence of sediment-laden ice. **However**, sediment concentrations were an order of magnitude less than samples from the ice canopy of 1978. The difference in load is related to a relatively rapid and quiet freeze-up, which resulted in reduced quantities of **frazil** ice and sediment. The lack of major storms and the presence of an ice cover on much of the inner shelf reduced sediment concentrations. The sediments present in the 1980-81 were concentrated inshore, at coastal promontories, and in shallow lagoons. These are the locations where downdrift accumulations of **frazil** ice were most likely to occur under the prevailing northeasterly.

REFERENCES

- Barnes, P. W., Reimnitz, Erk, and Fox, Dennis, 1982, Ice-rafting of **fine-grained** sediment, a sorting and transport mechanism, Beaufort Sea, Alaska: *Journal of Sedimentary Petrology*, v. 52, no. 2 (in press).
- Drake, D. E., **Kolpack**, R. L. , and Fischer, P. J., 1972, Sedimentary transport on the Santa Barbara-Oxnard Shelf, Santa Barbara Channel, California, in Swift, D. P. , Duane, D. B., and **Pilkey**, O. H. Eds., *Shelf Sediment Transport: Process and Pattern*, **Stroudsburg**, Pennsylvania, **Dowden Hutchinson**, and Ross, p. **307-332**.
- Reimnitz, Erk, and Dunton, K.H., **1979**, Diving observations on the soft-ice layer under the fast ice at **DS-11** in Stefansson Sound boulder patch: *Environmental Assessment of the Alaskan Continental Shelf, Annual Reports of Principal Investigators, March, 1979*, **NOAA/OCSEAP**, Boulder, **Colo.** , v. 9, p. 210-230s

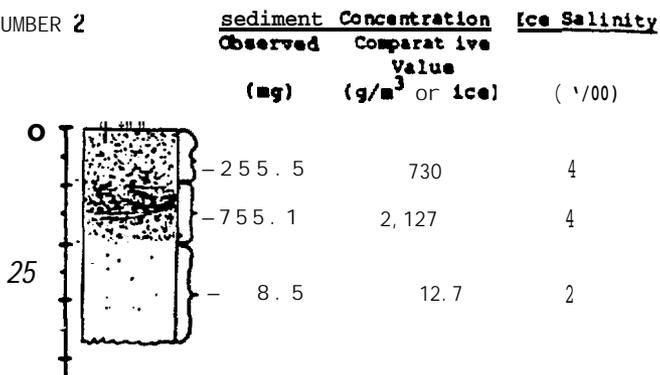
CORE NUMBER 2

Core Description

0-14 cm - gradually increasing sediment content with opaque ice at 0-2 cm and maximum sediment zone at 10-14 cm.

14-18 cm - intermediate amount of sediments with a jagged interface with clear ice at 18 cm.

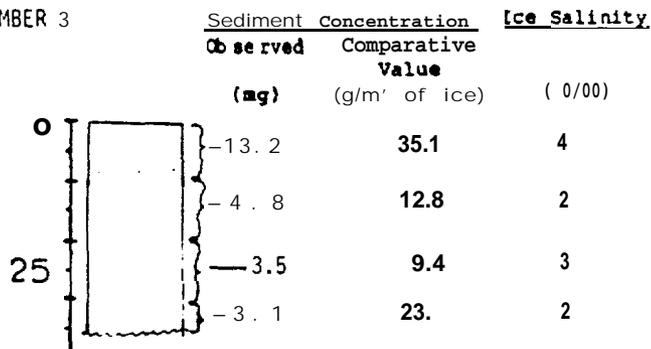
18-37.5 cm - clear ice



CORE NUMBER 3

Core description

0-35.5 cm - sediment-free core except for approximately 1-cm band at 8 cm of very light sedimentation. Opaque ice throughout.



CORE NUMBER 4

Core Description

0-1 cm - Several 2-cm-long detritus.

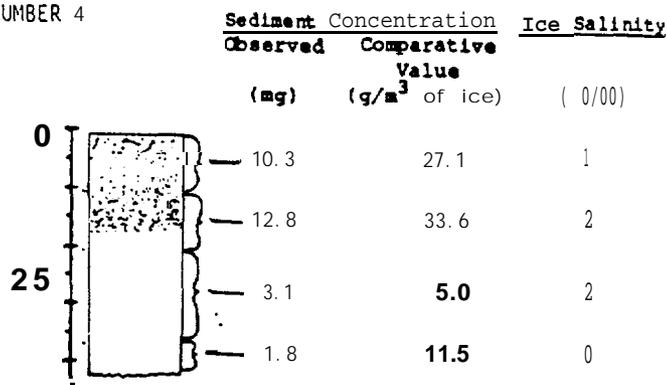
0-4.5 cm - Lightly sedimented section, obscured boundary with clean ice at 4.5 cm.

4.5-17 cm - Opaque ice at 4.5 cm gradually increasing sediments to an abrupt interface at 17 cm.

17-20 cm - Clear ice.

20-37 cm - Clear ice.

34-42 cm - Soft ice.



CORE NUMBER 5

Core Description

0-18 cm - Opaque ice - no layering or inclusions.

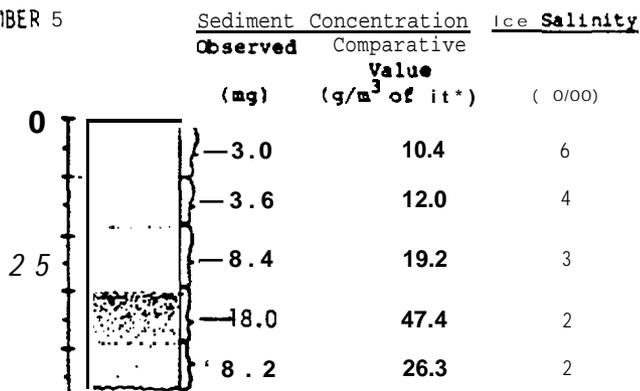
18-19 cm - Very faint sediment line - smooth boundaries.

19-30 cm - Relatively cleaner ice.

30-31 cm - Sediment layer - abrupt boundary at 30 cm - gradually decreasing sediments into next section. Uneven bottom boundary.

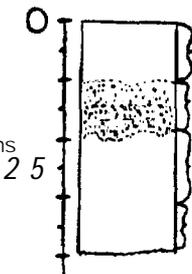
31-38.5 cm - Gradually decreasing sediment content - no inclusions.

38.5-47.5 cm - Dark plates.



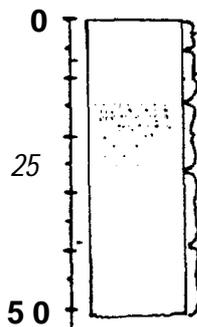
CORE NUMBER b

<u>Core Description</u>	<u>Sediment Concentration</u>		<u>Ice Salinity</u>
	<u>Observed</u> (mg)	<u>Comparative Value</u> (g/m ³ of ice)	(‰)
<u>0-6 cm</u> - Opaque ice with 3 distinct layers of patchy, cloudy ice. Densities are at 0-2 cm - 3-4 cm and 5-6 cm.	2.0	5.5	2
<u>6-10.5 cm</u> - Clear ice, no layering or inclusions	105.7	341.0	3
<u>10.5-14 cm</u> - Sediment-laden ice - uneven boundary at 19 cm.	2.9	5.9	2
<u>19-32 cm</u> - Opaque ice - no layering or inclusions.	3.1	14.0	2
<u>32-39.25 cm</u> - Soft, white, opaque ice - crystals prevalent.			



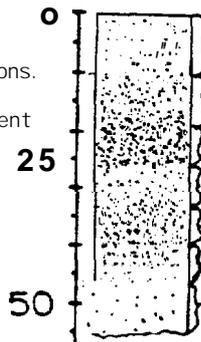
CORE NUMBER 7 "

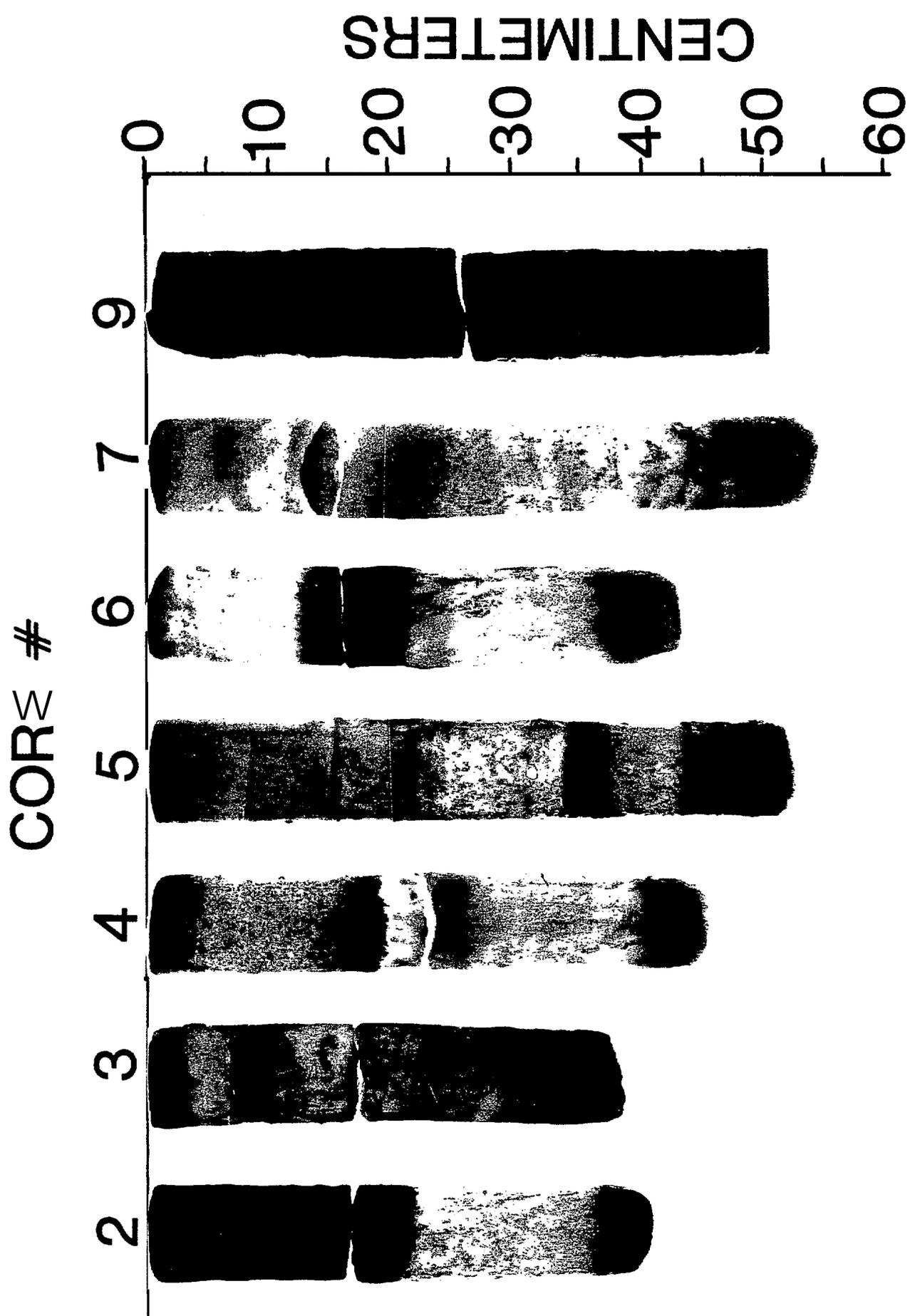
<u>Core Description</u>	<u>Sediment Concentration</u>		<u>Ice Salinity</u>
	<u>Observed</u> (mg)	<u>Comparative Value</u> (g/m ³ of ice)	(‰)
<u>0-15 cm</u> - Opaque cloudy ice - air bubbles - cleaner ice at bottom of section.	4.7	19.4	4
	3.2	14.4	3
<u>15-26 cm</u> - Sediment layer - denser at 15 cm, gradually decreasing concentrations to a non-distinct boundary at 26 cm.	18.0	40.9	2
	6.6	14.4	3
<u>26-38 cm</u> - Clear ice - no inclusions	8.0	19.6	2
<u>38-51 cm</u> - White, ice crystals, air bubbles, inclusions.			



CORE NUMBER B

<u>Core Description</u>	<u>Sediment Concentration</u>		<u>Ice Salinity</u>
	<u>Observed</u> (mg)	<u>Comparative Value</u> (g/m ³ of ice)	(‰)
<u>0-10 cm</u> - Dirty ice - some sediment inclusions.	21.0	56.8	0
<u>10-18.5 cm</u> - Very dirty ice - sediment inclusions.	86.3	274.0	2
<u>18.5-26.5 cm</u> - Dirty ice - increasing in sediment concentration towards bottom - any sediment inclusions.	48.6	160.4	0
	32.4	118.}	2
<u>26.5-28 cm</u> - Diagonal sediment band	52.2	193.4	2
	26.6	101.1	0
<u>28-33 cm</u> - Dirty ice - no sediment inclusions.	21.7	90.0	0
<u>33-47 cm</u> - Dirty ice - sediment inclusions.			
<u>47-55 cm</u> - Opaque ice, large crystals.			





CORE NUMBER 9

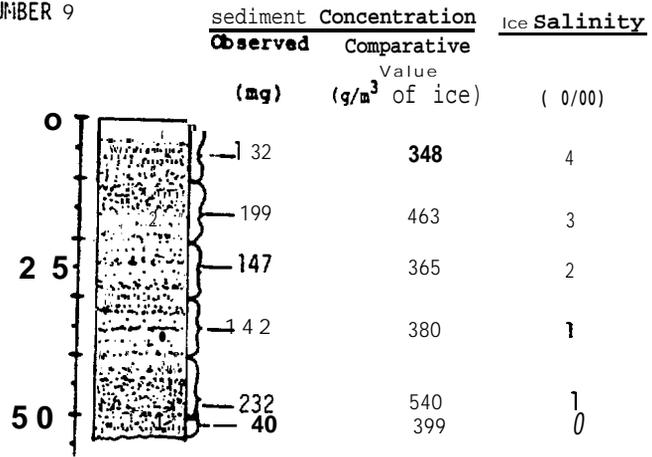
Core Description

0-3 cm - "Clean" ice.

3-9 cm - Dirty ice, abrupt contact - irregular surface at 5 cm.

9-16 cm - Increasing sediment content

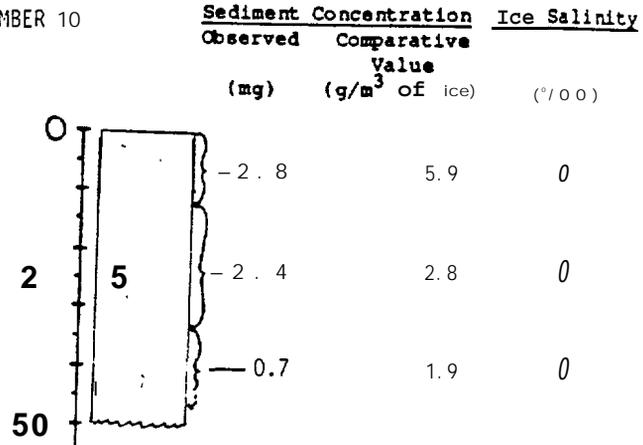
16-44 cm - Abrupt but ragged contact with clearer ice at 16 cm with faint turbidity bands at 28-34 cm.



CORE NUMBER 10

Core Description

0-46 cm - Clear ice - freshwater ice?



CORE NUMBER 11

Core Description

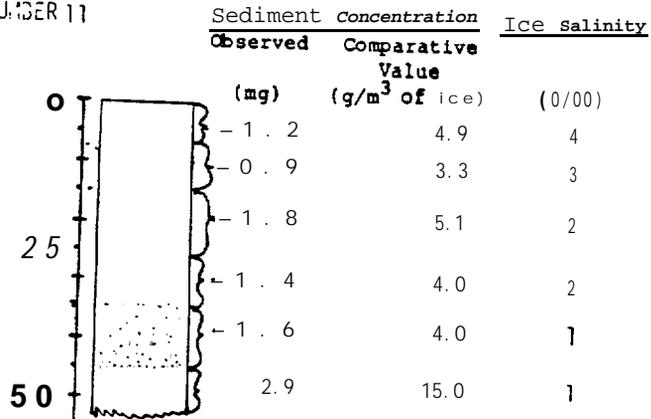
0-7.5 cm - Opaque ice.

7.5-15.5 cm - Very clear ice.

15.5-35 cm - Opaque ice.

35-46 cm - Gradually increasing sediment contact - cloudy-opaque.

46-53 cm - Ice crystals - uneven boundary surfaces.



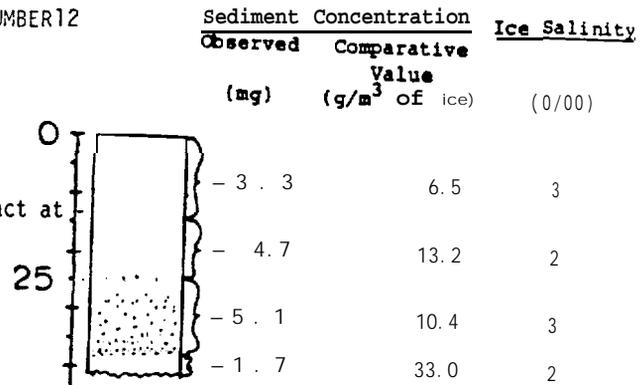
CORE NUMBER 12

Core Description

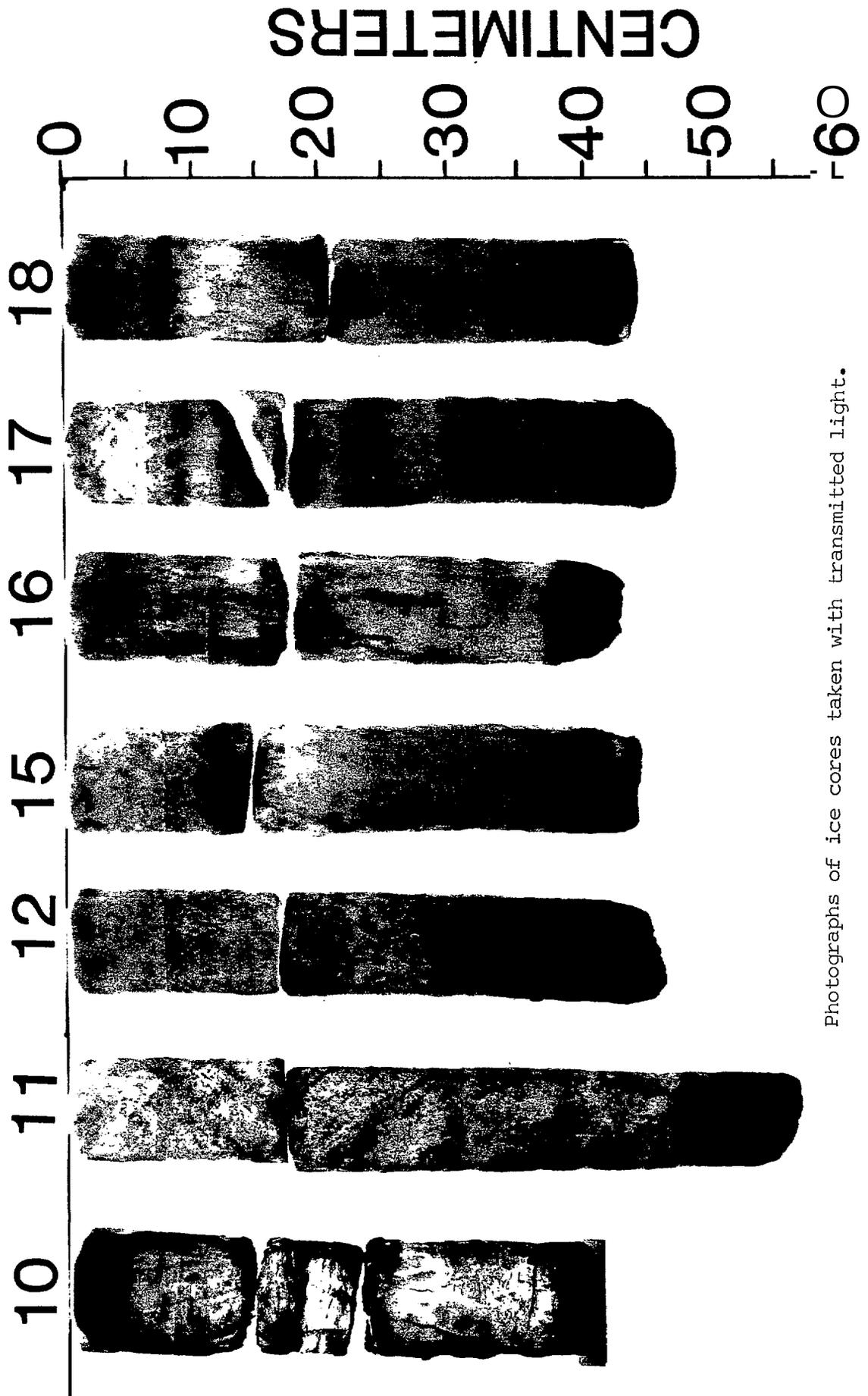
0-25 cm - Opaque ice, no visible sediments.

25-38 cm - Gradually increasing sediment content obscured interface at 25 cm - abrupt contact at 38 cm.

38-41 cm - Distinct ice crystals arranged longitudinally.



CORE #



Photographs of ice cores taken with transmitted light.

CORE NUMBER 13

Core Description

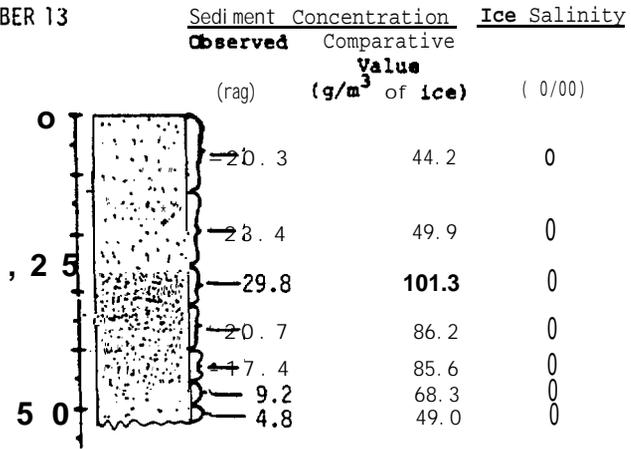
0-13.5 cm - Opaque ice, uniform light sediment distribution.

13.5-26 cm - Uniform moderate sediment distribution.

26-33.5 cm - Turbid ice - sediment inclusions throughout.

33.5-40 cm - Light sediment inclusions.

40-45 cm - Moderate quantity of sedimentary inclusions throughout.



CORE NUMBER 14

Core Description

0-3 cm - Clear ice, uneven interface at 3 cm.

3-5.5 cm - Clear ice - freshwater? diagonal interface at 5 to 6 cm.

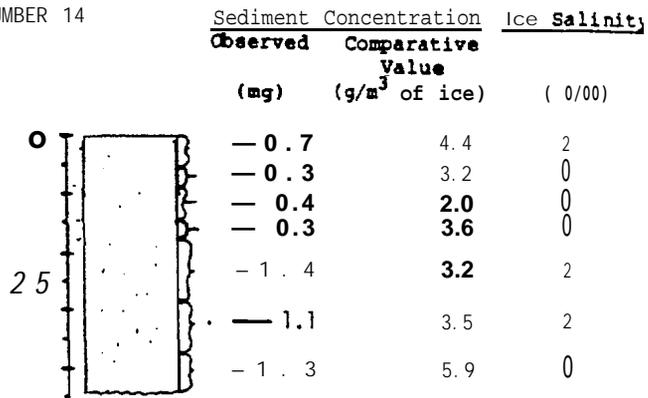
5.5-12.5 cm - Cloudy ice - no sediment. Distinct interface at 12.5 cm.

12.5-14.5 cm - Clear (freshwater?) ice

14.5-17.5 cm - Slightly opaque cloudy ice - abrupt interface at 17.5 cm.

17.5-37 cm - Cloudy ice with uniform sediment coloration.

37-44.5 cm - Ice crystals with many air inclusions.



CORE NUMBER 15

Core Description

0-4 cm - Opaque ice.

4-11 cm - Increasing sediment content - no distinct inclusions

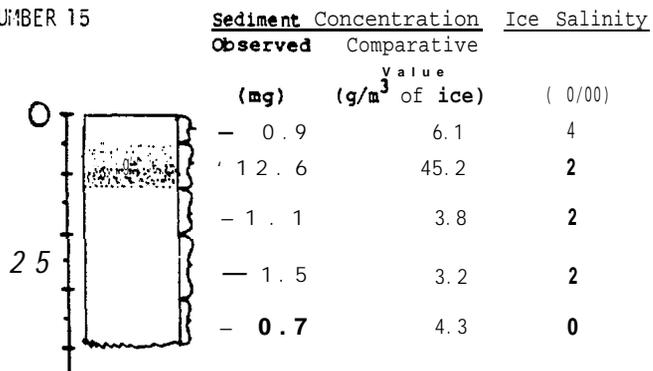
10-11 cm - Uneven band of sediment - very distinct interface at 11 cm

11-12 cm - Opaque ice - no sediment visible.

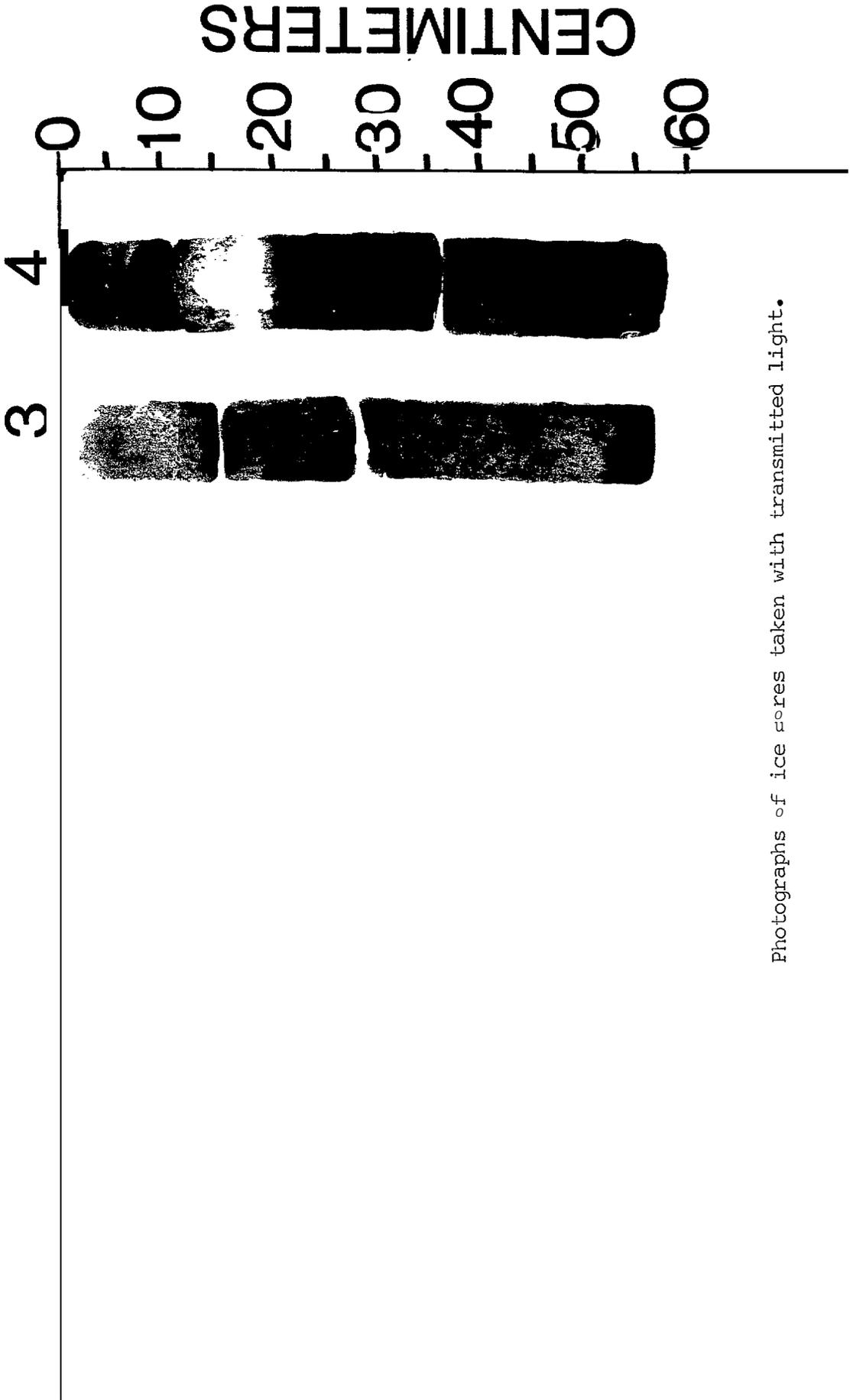
12-20 cm - Relatively clear ice, no sediment.

20-32.5 cm - Gradually-increasing sediment content.

32.5-38.8 cm - Uneven longitudinal crystals-patches of very clear ice.



CORE



Photographs of ice cores taken with transmitted light.

CORE NUMBER 16

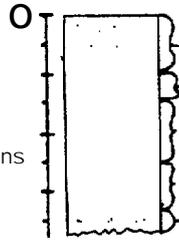
Core Description

0-9 cm - Opaque ice - no distinct sediment inclusions.

9-14.8 cm - Clearer ice.

14.8-33.5 cm - Opaque ice, sediment concentrations decreases with core depth - no sediment bands or layers.

33.5-37.5 cm - Ice crystals, no sediment inclusions.



<u>Sediment Observed</u>	<u>Concentration Comparative Value</u>	<u>Ice Salinity</u>
(mg)	(g/m ³ of ice)	(0/00)

- 2.1	6.1	5
- 0.6	3.3	2
- 0.8	2.0	0
- 0.4	1.4	0
- 0.5	4.8	2

CORE NUMBER 17

Core Description

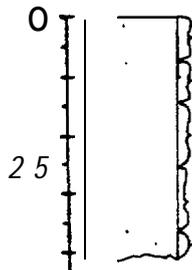
0-3.5 cm - Cloudy ice - no distinct sediment.

3.5-7 cm - Clear ice.

7-7.5 cm - Cloudy band - no distinct sediment.

7.5-10 cm - Clear ice.

10-11 cm - Cloudy ice - no distinct sediment.



<u>Sediment Observed</u>	<u>Concentration Comparative Value</u>	<u>Ice Salinity</u>
(mg)	(g/m ³ of ice)	(0/00)

- 1.1	4.2	3
- 1.0	4.3	2
- 0.9	2.1	2
- 0.8	2.3	2
- 0.8	5.5	1

CORE NUMBER 18

Core Description

0-6.5 cm - Cloudy ice - with sediment - no distinct sediment accumulations.

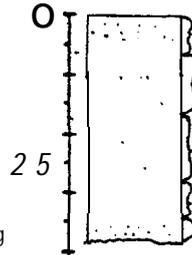
6.5-7 cm - Dark band - no distinct sediment.

7-17.3 cm - Clear ice gradually getting darker - no sediment - sharp interface at 7cm.

17.3-28.4 cm - Opaque ice - gradually increasing in cloudiness - no sediment inclusions.

28.4-35 cm - Uniformly spaced sediment bands approximately 4 cm apart. Sediment concentration- increasing towards bottom.

35-38.5 cm - Ice crystals - heaviest sediment concentration (darkest area).



<u>Sediment Observed</u>	<u>Concentration Comparative Value</u>	<u>Ice salinity</u>
(mg)	(g/m ³ of ice)	(0/00)

- 3.9	16.2	5
- 2.9	6.7	2
- 1.6	3.8	2
- 1.2	4.6	2
- 0.8	8.9	2